
SEAMOCs Workshop
Runup of the sea waves on a beach
Ira Didenkulova
Institute of Applied Physics, Russian Academy of Sciences, Nizhny Novgorod, Russia
Institute of Cybernetics, Tallinn, Estonia

Tsunami Wave Shapes at Japanese Coast
26 May 1983
Japan Sea
(Shuto, 1983)

Nonlinear Shallow Water Theory

$$\frac{\partial \eta}{\partial t} + \frac{\partial}{\partial x} [(-\alpha x + \eta)u] = 0$$

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + g \frac{\partial \eta}{\partial x} = 0$$

$h(x) = -\alpha x$

Water waves of finite amplitude on a sloping beach

By G. F. CARRIER and H. P. GREENSPAN
Pierce Hall, Harvard University
 (Received 2 December 1957)

SUMMARY

In this paper, we investigate the behaviour of a wave as it climbs a sloping beach. Explicit solutions of the equations of the non-linear inviscid shallow-water theory are obtained for several physically interesting wave-forms. In particular it is shown that waves can climb a sloping beach without breaking. Formulae for the motions of the instantaneous shoreline as well as the time histories of specific wave-forms are presented.

Hodograph Transformation

$$\frac{\partial^2 \Phi}{\partial \lambda^2} - \frac{\partial^2 \Phi}{\partial \sigma^2} - \frac{1}{\sigma} \frac{\partial \Phi}{\partial \sigma} = 0$$

$$\sigma = 2\sqrt{g(h + \eta)} \geq 0$$

Implicit form

$$\eta = \frac{1}{2g} \left(\frac{\partial \Phi}{\partial \lambda} - u^2 \right)$$

$$x = \frac{1}{2\alpha g} \left(\frac{\partial \Phi}{\partial \lambda} - u^2 - \frac{\sigma^2}{2} \right)$$

$$u = \frac{1}{\sigma} \frac{\partial \Phi}{\partial \sigma}$$

$$t = \frac{1}{\alpha g} \left(\lambda - \frac{1}{\sigma} \frac{\partial \Phi}{\partial \sigma} \right)$$

Implicit Form – A Few Exact and Approximated Solutions For Whole Field as Function of (x,t)

Papers (after 2000):

Carrier G.F., Wu T.T., Yeh H. Tsunami run-up and draw-down on a plane beach. *J. Fluid Mech.*, 2003, v. 475, 79-99.

Kanoğlu U. Nonlinear evolution and runup-rundown of long waves over a sloping beach. *J. Fluid Mech.*, 2004, v. 513, 363-372.

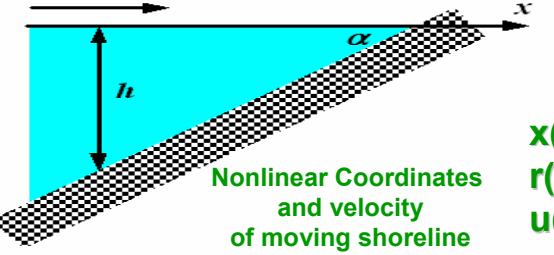
Tinti S., Tonini R. Analytical evolution of tsunamis induced by near-shore earthquakes on a constant-slope ocean. *J. Fluid Mech.*, 2005, v. 535, 33-64.

Kanoğlu U., and Synolakis, C. Initial value problem solution of nonlinear shallow water-wave equations. *Phys. Review Lett.*, 2006, v. 97, 148501.

Explicit Solution for Moving Shoreline
 if Incident Wave is given Far from Shoreline
 where It is Linear

Pelinovsky & Mazova, 1992

$$u(t) = U\left(t + \frac{u}{\alpha g}\right) \quad r(t) = \alpha \int u(t) dt$$



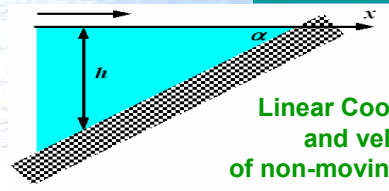
Nonlinear Coordinates and velocity of moving shoreline
 $x(t)$
 $r(t)$
 $u(t)$

First Step – Solution of Linear Equations
 For Wave Transformation on Beach

Incident Wave
 $\eta(t) = \sum_{n=1}^{\infty} A_n \sin(n\omega t + \varphi_n)$

$$R(t) = 2\pi \sqrt{\frac{2L}{\lambda}} \sum_{n=1}^{\infty} \sqrt{n} A_n \sin\left(n\omega t + \varphi_n + \frac{\pi}{4}\right)$$

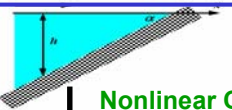
$$U(t) = \frac{1}{\alpha} \frac{dR}{dt}$$



Linear Coordinates and velocity of non-moving shoreline
 $x=0$
 $R(t)$
 $U(t)$

Second Step – “Nonlinear” Moving Shoreline

$$u(t) = U\left(t + \frac{u}{\alpha g}\right) \quad r(t) = \alpha \int u(t) dt$$



Linear Coordinates and velocity of non-moving shoreline
 $x=0, R(t), U(t)$

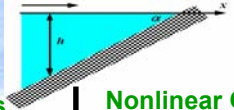
Nonlinear Coordinates and velocity of moving shoreline
 $x(t), r(t), u(t)$

First Result: Linear Theory predicts Maxima

$$u_{\max} = U_{\max}$$

$$r_{\max} = R_{\max}$$

for sine wave
 $\frac{R}{H_0} = 2\pi \sqrt{\frac{2L}{\lambda_0}}$



Linear Coordinates and velocity of non-moving shoreline
 $x=0, R(t), U(t)$

Nonlinear Coordinates and velocity of moving shoreline
 $x(t), r(t), u(t)$

Natural Hazards 4: 221–234, 1991.
 © 1991 Kluwer Academic Publishers. Printed in the Netherlands.

**Tsunami Runup on Steep Slopes:
 How Good Linear Theory Really Is**

COSTAS EMMANUEL SYNOLAKIS
 School of Engineering, University of Southern California, Los Angeles, CA 90089-2531, U.S.A.

(Received: 28 January 1990; revised: 29 June 1990)

Abstract. This is a study of the application of linear theory for the estimation of the maximum runup height of long waves on plane beaches. The linear theory is reviewed and a method is presented for calculating the maximum runup. This method involves the calculation of the maximum value of an integral, now known as the runup integral. Laboratory and numerical results are presented to support this method. The implications of the theory are used to reevaluate many existing empirical runup correlations. It is shown that linear theory predicts the maximum runup satisfactorily. This study demonstrates that it is now possible to match complex offshore wave-evolution algorithms with linear theory runup solutions for the purpose of obtaining realistic tsunami inundation estimates.

Second Result: Linear Theory “predicts” Wave Breaking

$$u(t) = U\left(t + \frac{u}{\alpha g}\right)$$

$$\frac{\partial u}{\partial t} = \frac{U'}{1 - \frac{U'}{\alpha g}}$$

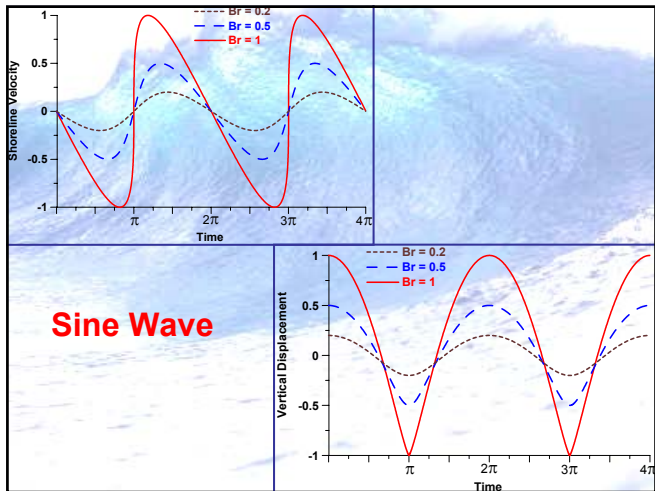
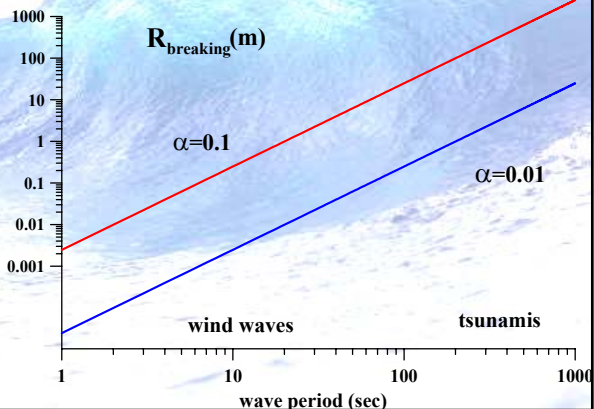
$$Br = \frac{1}{g\alpha^2} \max\left(\frac{d^2 R}{dt^2}\right) = 1$$

The same as Jacobian Transformation

$$\frac{\partial(t, x)}{\partial(I_+, I_-)} = 0$$

“Linear” Vertical Acceleration = Gravity Acceleration

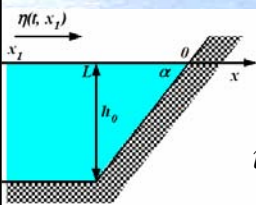
75% of tsunamis are not breaking waves,
Mazova, Pelinovsky and Soloviev, 1982



Nonlinear Deformation

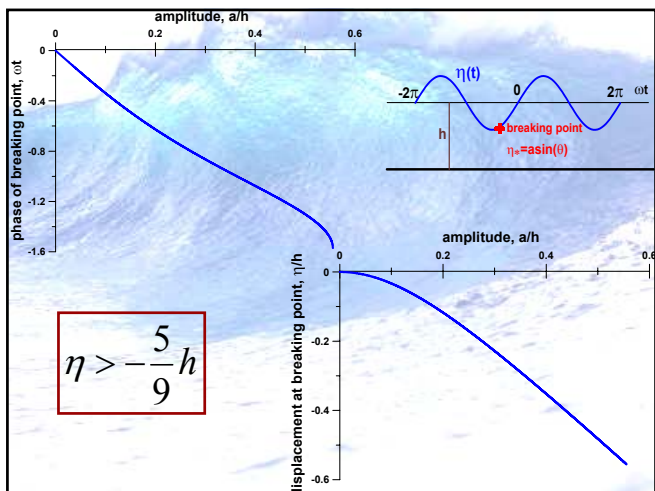
$$\frac{\partial \eta}{\partial t} + V \frac{\partial \eta}{\partial x} = 0$$

$$\eta(x, t) = \eta_0 \left(t - \frac{x}{V(\eta)} \right)$$

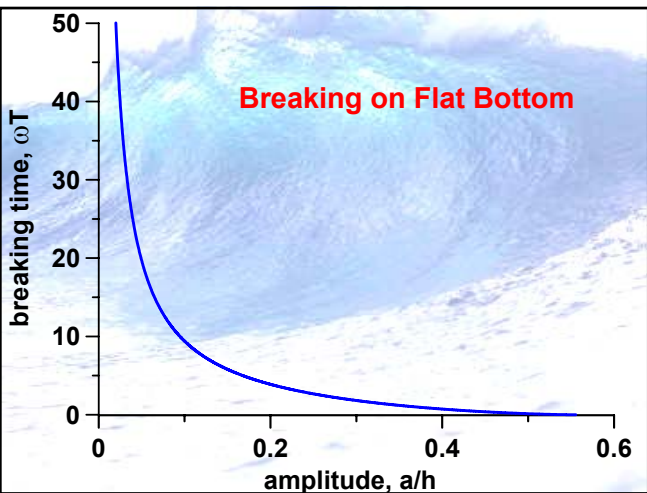


$$V = 3\sqrt{g(h + \eta)} - 2\sqrt{gh}$$

$$u = 2(\sqrt{g(h + \eta)} - \sqrt{gh})$$



$$\eta > -\frac{5}{9}h$$



Indian Ocean Estimates

$$X \sim 0.1 \lambda \frac{h}{a}$$

$$X = \frac{2\sqrt{gh}h}{3\omega a} = \frac{\lambda h}{3\pi a}$$

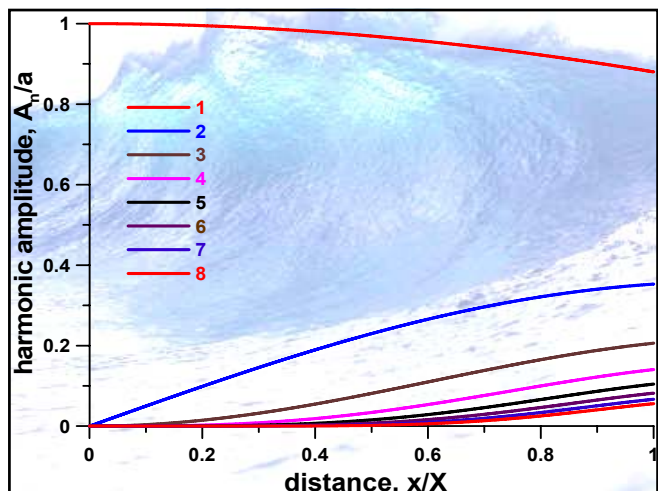
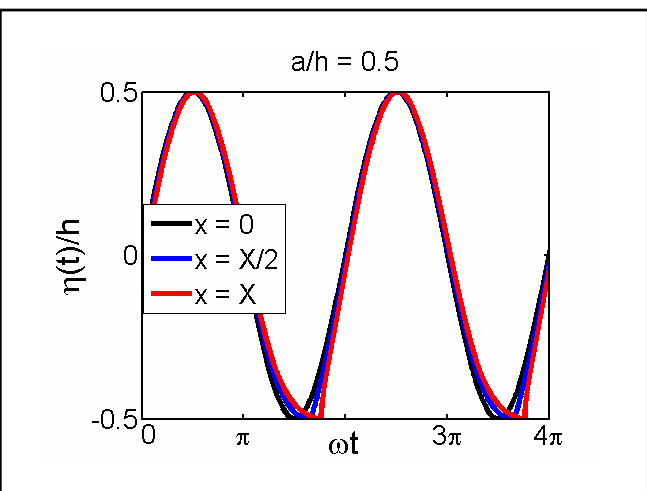
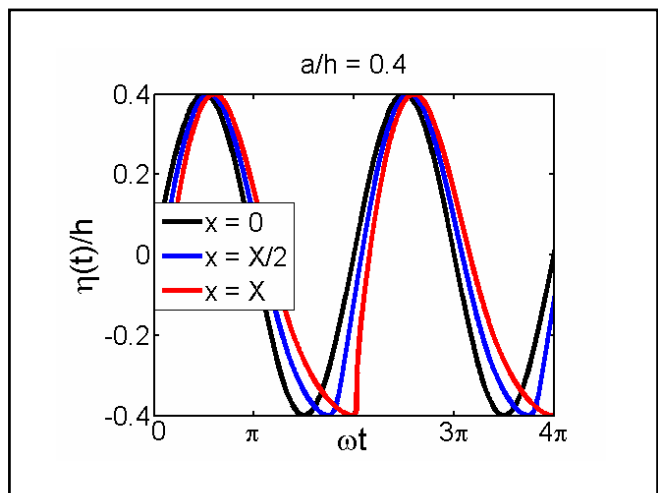
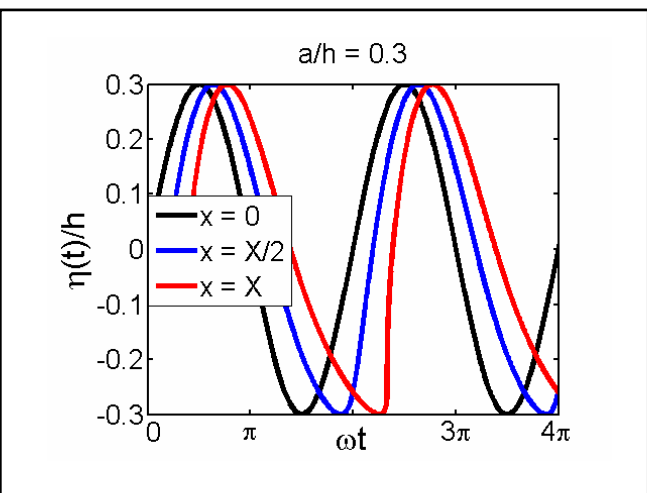
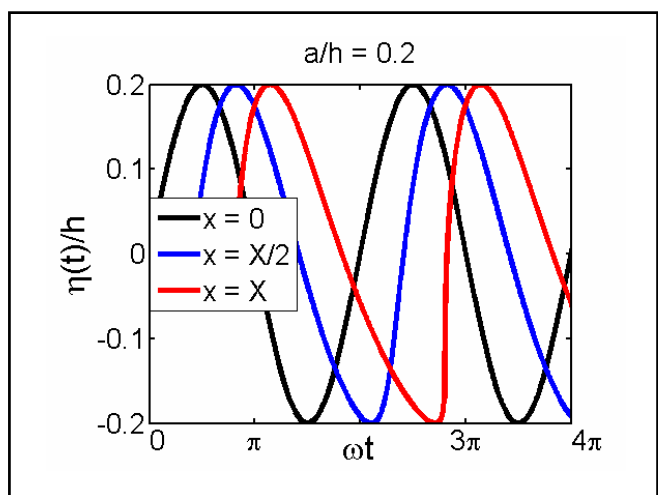
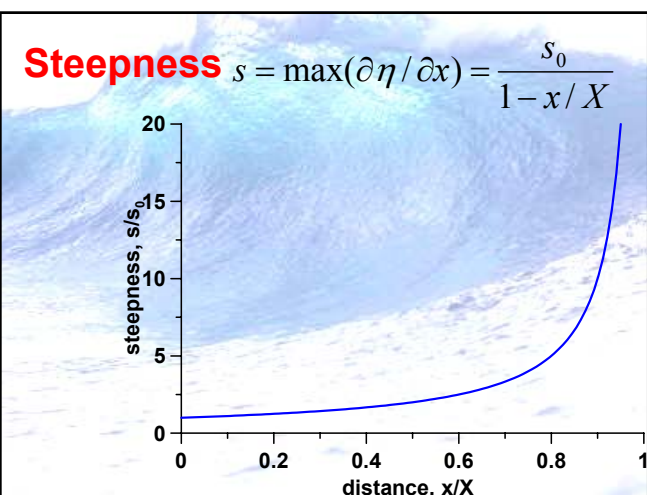
$h = 1 \text{ km}$

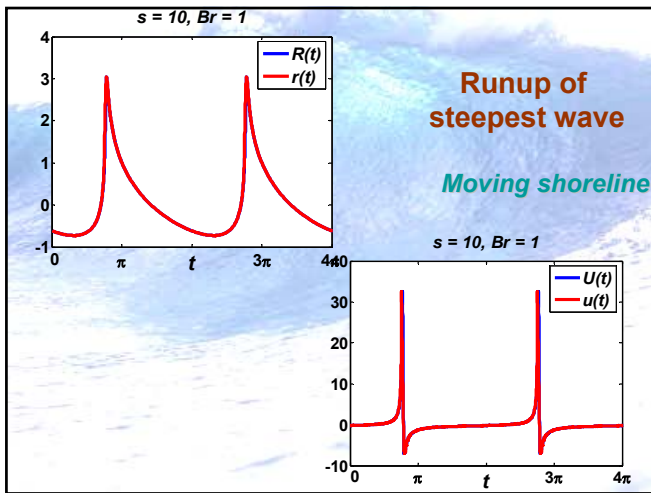
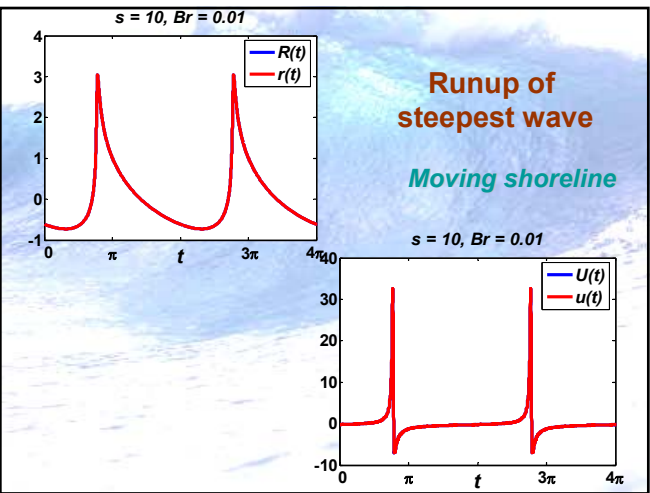
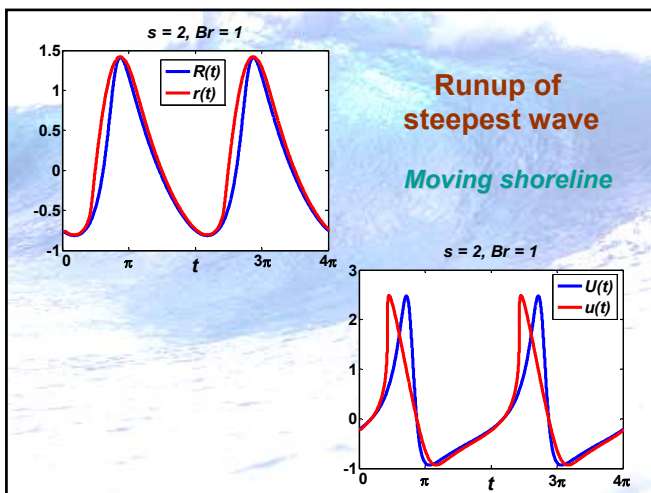
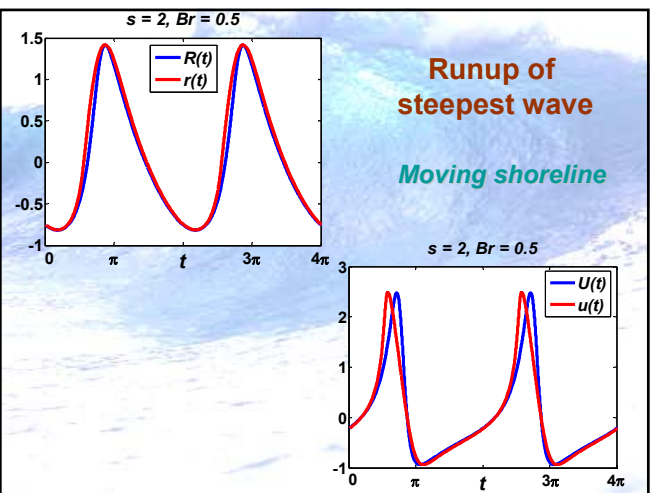
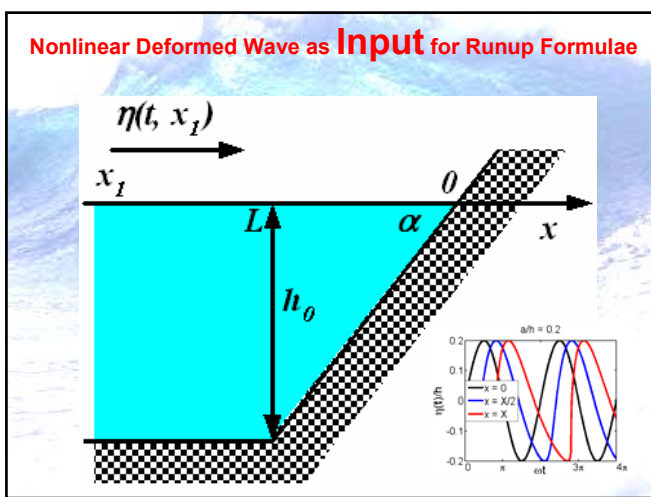
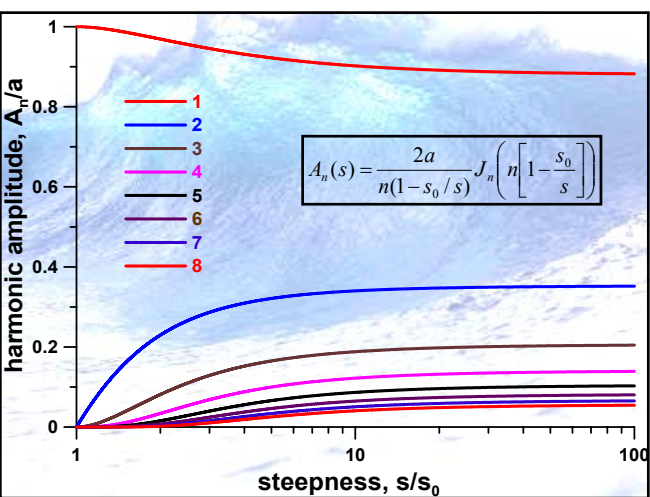
$X \sim 2500 \text{ km}$

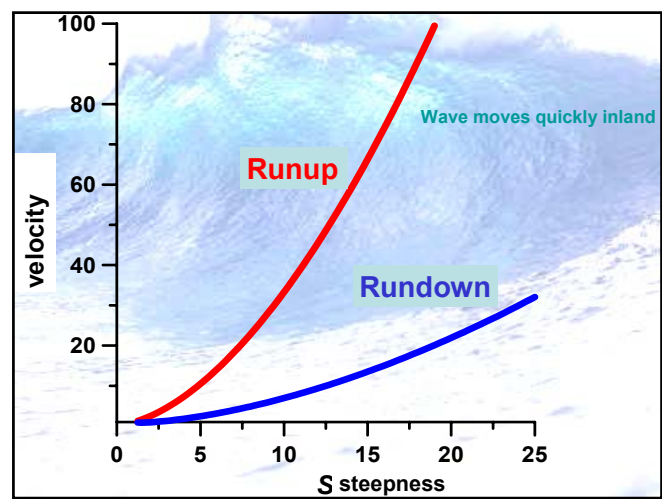
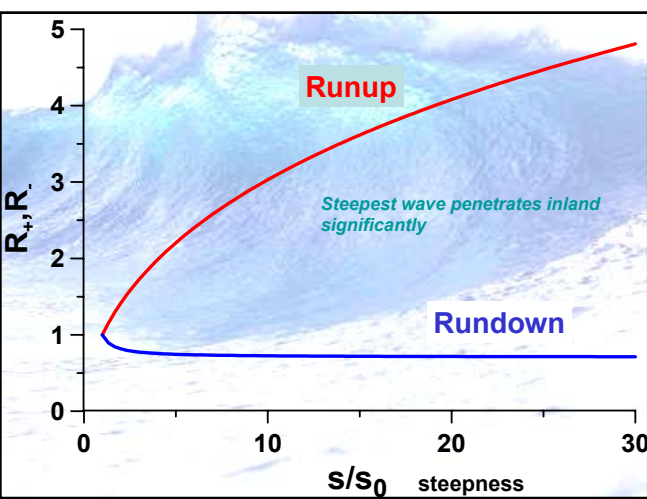
$a = 2 \text{ m}$

$\lambda = 50 \text{ km}$

To India!







ISSN 1025-3340, Doklady Earth Sciences, 2006, Vol. 411, No. 5, pp. 1241-1243. © Pleiades Publishing, Inc., 2006.
Original Russian Text © I.I. Didenkulova, N. Zahibo, A.A. Kurkin, E.N. Pelinovsky, T. Soomere, 2006, published in Doklady Akademii Nauk, 2006, Vol. 410, No. 5, pp. 670-676.

G E O P H Y S I C S

**Runup of Nonlinearly Deformed Waves
on a Coast**

I. I. Didenkulova^{a, b}, N. Zahibo^c, A. A. Kurkin^b, Corresponding Member of the RAS B. V. Levin^{d, e},
E. N. Pelinovsky^{a, b}, and T. Soomere^f
Received February 2, 2006

ISSN 0001-4338, Izvestiya, Atmospheric and Oceanic Physics, 2006, Vol. 42, No. 6, pp. 773-776. © Pleiades Publishing, Inc., 2006.
Original Russian Text © I.I. Didenkulova, N. Zahibo, A.A. Kurkin, E.N. Pelinovsky, 2006, published in Izvestiya AN. Fizika Atmosfery i Okeana, 2006, Vol. 42, No. 6, pp. 839-842.

**Steepness and Spectrum of a Nonlinearly Deformed Wave
on Shallow Waters**

I.I. Didenkulova^{a, b}, N. Zahibo^c, A. A. Kurkin^d, and E. N. Pelinovsky^{a, b}

**Does Runup Height depend on
the Incident Wave Shape if the
Incident Wave is symmetrical?**

For periodic wave – yes!

The anomalous behavior of the runup of cnoidal waves

Costas Emmanouel Synolakis and Manas Kumar Deb
School of Engineering, University of Southern California, Los Angeles, California 90089
James Eric Skjelbreia
MARINETEK, Hakonsongst 34, 7002 Trondheim, Norway

Phys. Fluids, 1988, v. 31, No. 1

For single wave???

Incident wave shapes used:

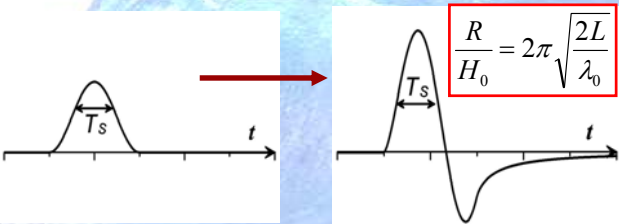
1. **Solitary Wave** $f(x) = \text{sech}^2(x)$
2. **Gaussian Pulse** $f(x) = \exp(-x^2)$
3. **Lorentz Pulse** $f(x) = \frac{1}{1+x^2}$

and several others

$$\frac{R}{H_0} = 2\pi \sqrt{\frac{2L}{\lambda_0}}$$

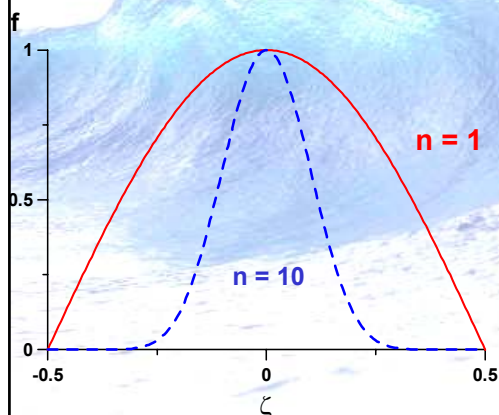
**How this formula
can be modified?**

Wave Length (Duration) Definition for Pulse

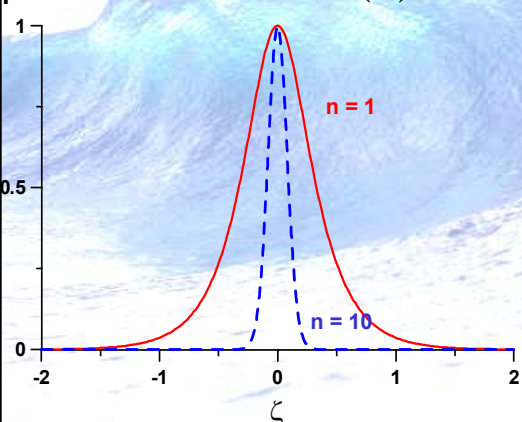


We suggest to use length of the wave on 2/3 level – philosophy used to define the significant wave properties

Incident Wave Shapes $f(x) = \cos^n(\pi x)$



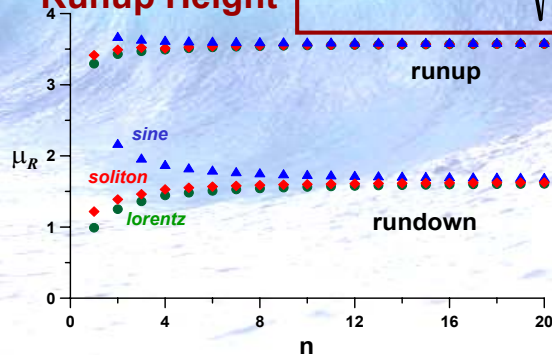
Incident Wave Shapes $f(x) = \text{sech}^n(4x)$



$R = 2\pi H_0 \sqrt{\frac{2L}{\lambda_0}}$

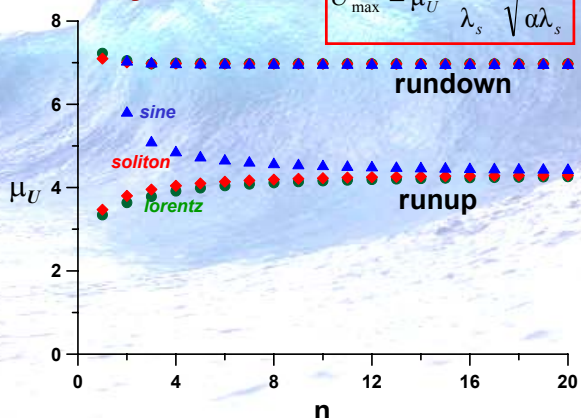
$R_{\max} = \mu_R^+ H_0 \sqrt{\frac{L}{\lambda_s}}$

Runup Height

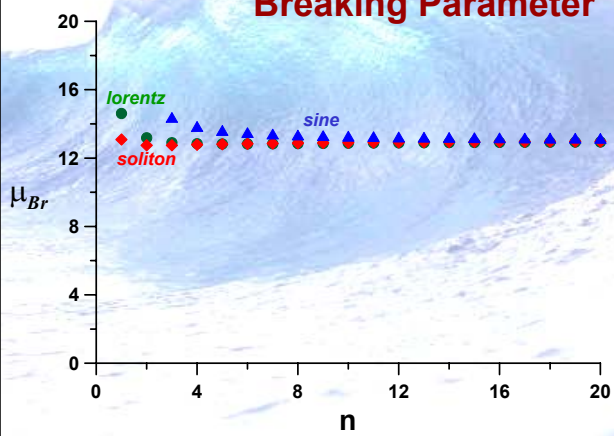


Velocity

$U_{\max} = \mu_U^+ \frac{H_0 L}{\lambda_s} \sqrt{\frac{g}{\alpha \lambda_s}}$



Breaking Parameter



Parametrized Formulas

$$R_{\text{runup}} = 3.5H_0 \sqrt{\frac{L}{\lambda_s}}$$

$$R_{\text{down}} = 1.5H_0 \sqrt{\frac{L}{\lambda_s}}$$

$$U_{\text{runup}} = 4.5 \frac{H_0 L}{\lambda_s} \sqrt{\frac{g}{\alpha \lambda_s}}$$

$$U_{\text{rundown}} = 7 \frac{H_0 L}{\lambda_s} \sqrt{\frac{g}{\alpha \lambda_s}}$$

$$Br = 13 \frac{H_0 L}{\alpha \lambda_s^2} \sqrt{\frac{L}{\lambda_s}}$$

Applications:

Easy Estimates of Runup Characteristics

ISSN 0001-4304, *Izvestiya, Atmospheric and Oceanic Physics*, 2007, Vol. 43, No. 3, pp. 384–390 © Pleiades Publishing, Ltd., 2007.
Original Russian Text © I.I. Didenkulova, A.A. Kurkin, E.N. Pelinovsky, 2007, published in *Izvestiya AN, Fizika Atmosfery i Okeana*, 2007, Vol. 43, No. 3, pp. 419–425.

Run-up of Solitary Waves on Slopes with Different Profiles

I. I. Didenkulova^{a,b}, A. A. Kurkin^a, and E. N. Pelinovsky^{a,b}

^a Nizhni Novgorod State Technical University, ul. Minina 24, Nizhni Novgorod, 603950 Russia
e-mail: kurkin@kis.ru

^b Institute of Applied Physics, Russian Academy of Sciences, ul. Ul'yanova 46, Nizhni Novgorod, 603950 Russia
Received September 7, 2006

Conclusion:

**Steepest Wave Penetrates Inland
on Large Distance and with Large Velocity
- and Slowly into the Sea**

**Formulas for Solitary Wave Runup
can be Parameterized**

**These results are important for express
estimates**