

Climate Change – The Physical Background

Andreas Sterl

KNMI

- Basics of the climate system
- Anthropogenic influence
- Projected changes & impact

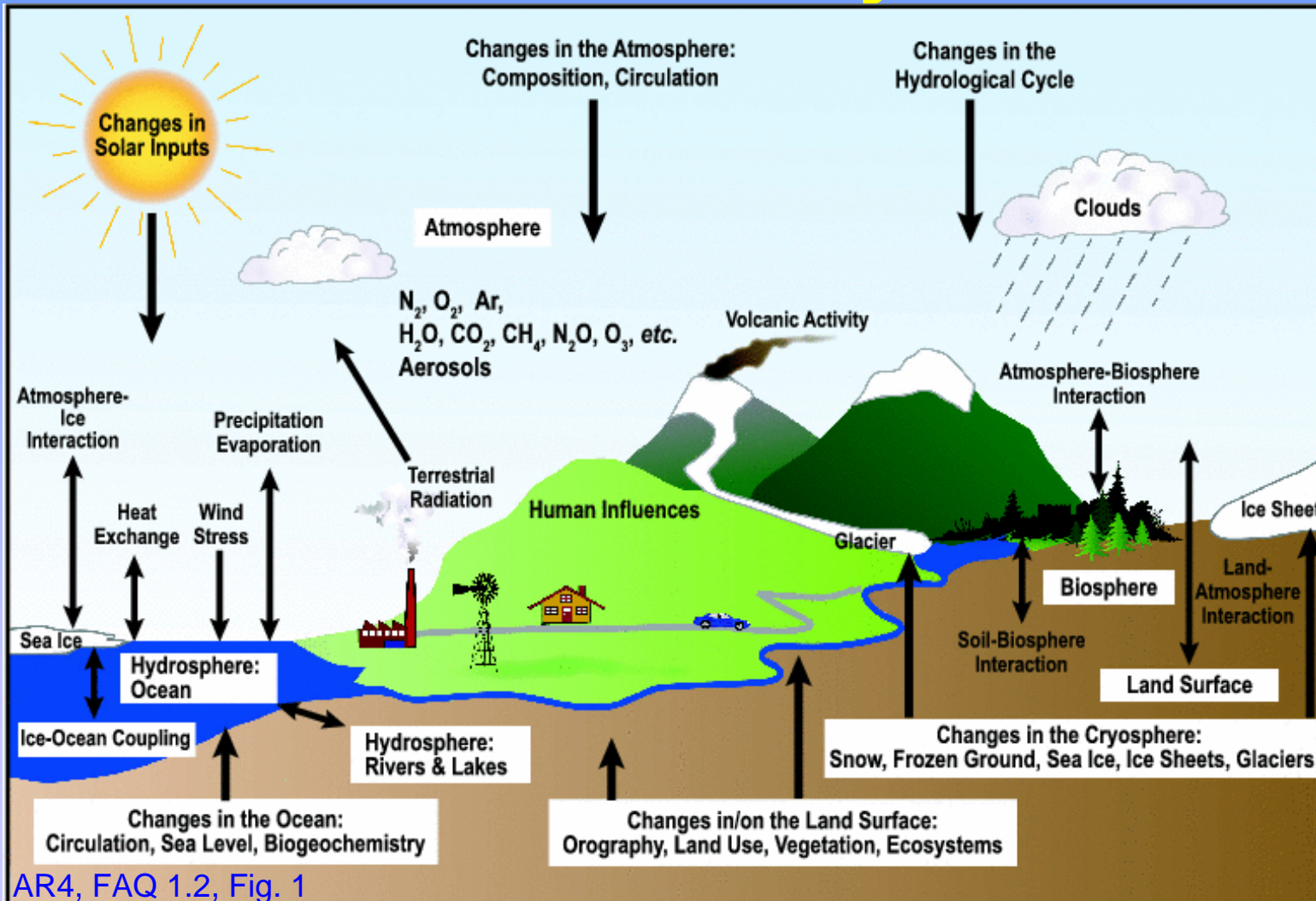
Climate – what's that?

Climate = statistics of weather

Weather = state of atmosphere
at a particular time

“Climate is what you expect,
weather is what you get”

The climate system

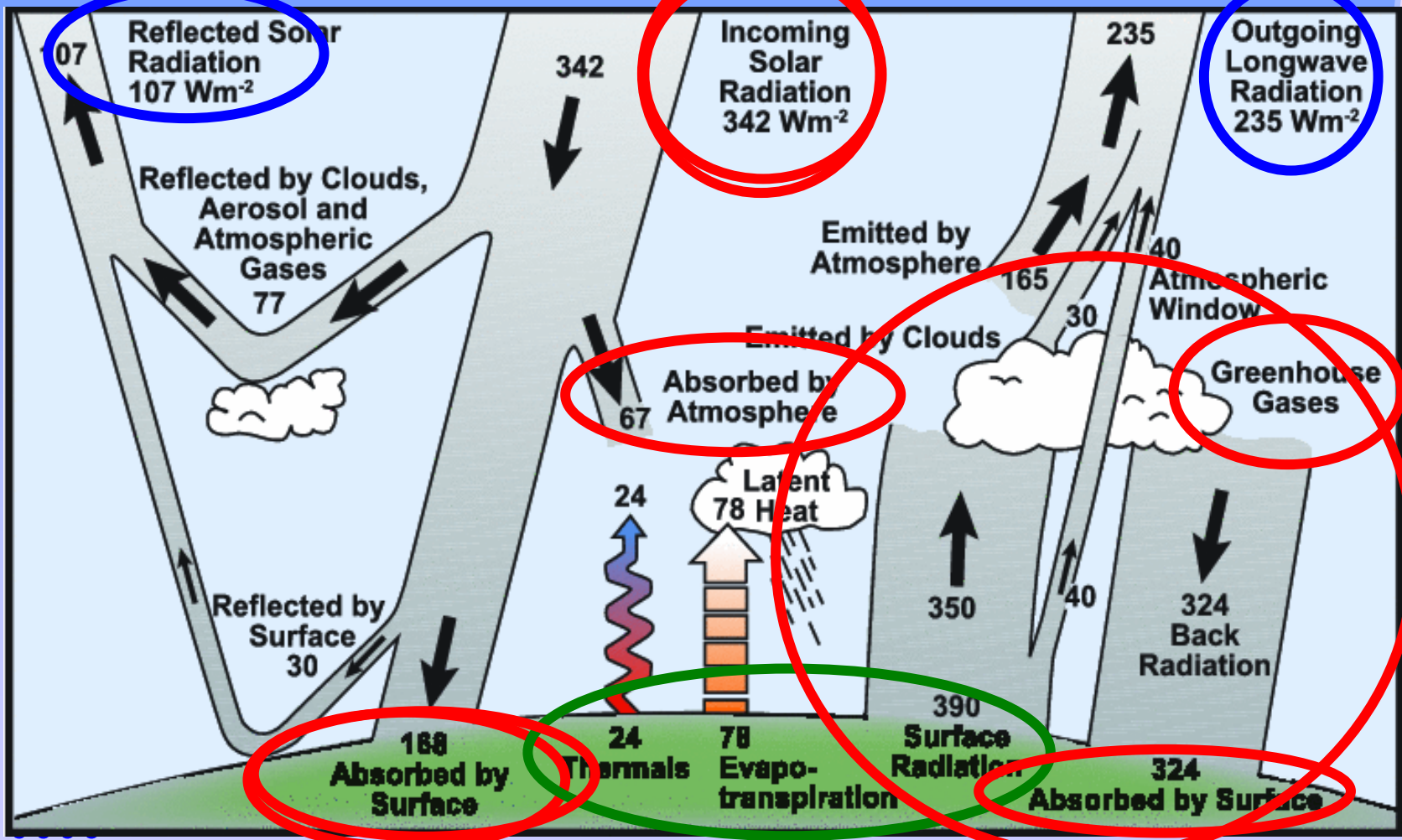


AR4, FAQ 1.2, Fig. 1

••• How does the climate system work?

- short-wave radiation (sun) heats (tropics)
- atmosphere and ocean transport heat poleward
- modulated by continents
- long-wave radiation into space cools

Radiative forcing



AR4, FAQ 1.1, Fig. 1

Andreas Sterl, SEAMOCS workshop, Palmse, 11.10.2007

Greenhouse effect

Outgoing longwave radiation is absorbed in the atmosphere and radiated back to the surface.

without: $T_{\text{glob}} = -15^{\circ}\text{C}$

observed: $T_{\text{glob}} = +15^{\circ}\text{C}$

Greenhouse gases

only 1 % of atmosphere

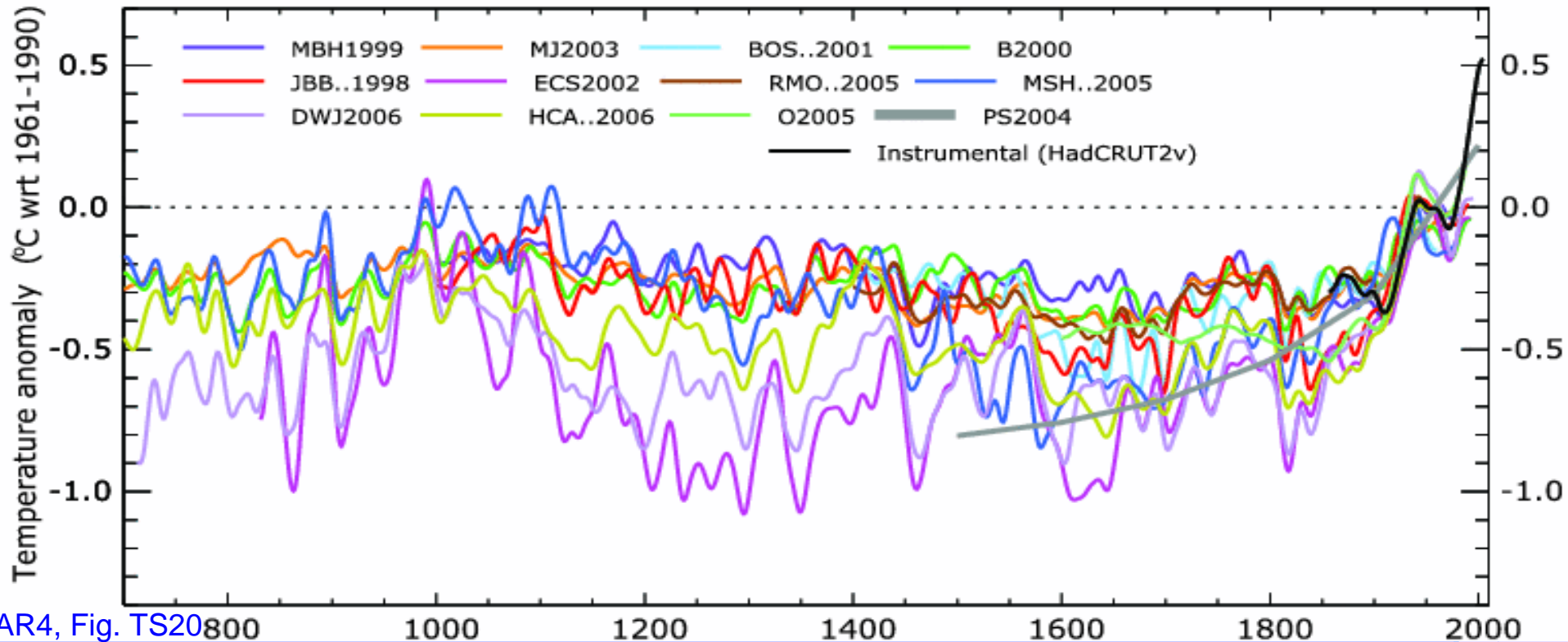
- H_2O 70 % <= feedback (~2x)
- CO_2 15 %
- CH_4)
- N_2O) 5 %
- Ozone, ...)
- (clouds 10 %) <= feedback (?)

(van Dorland, 1999, p. 17/18)

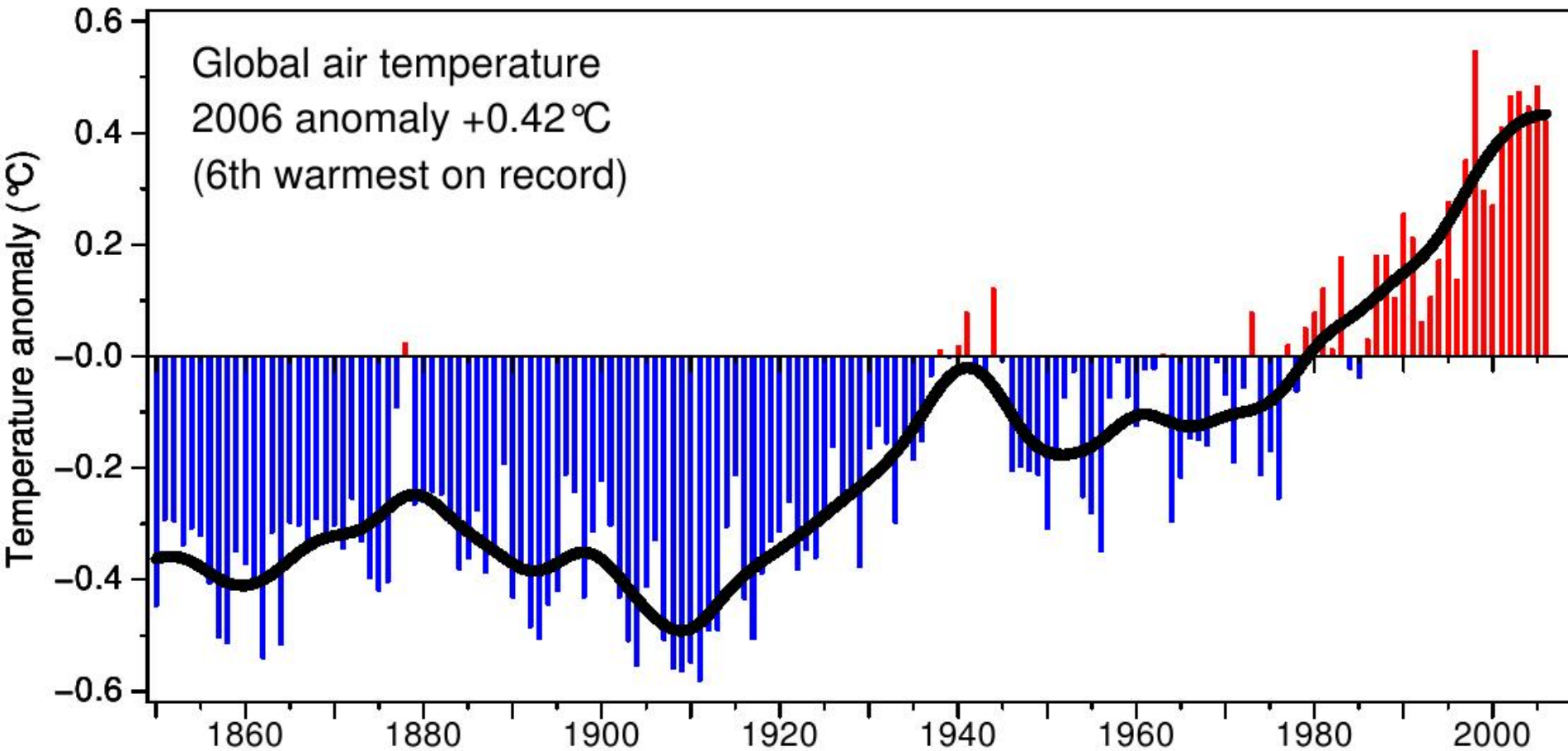
Observed climate change

The past 1300 years

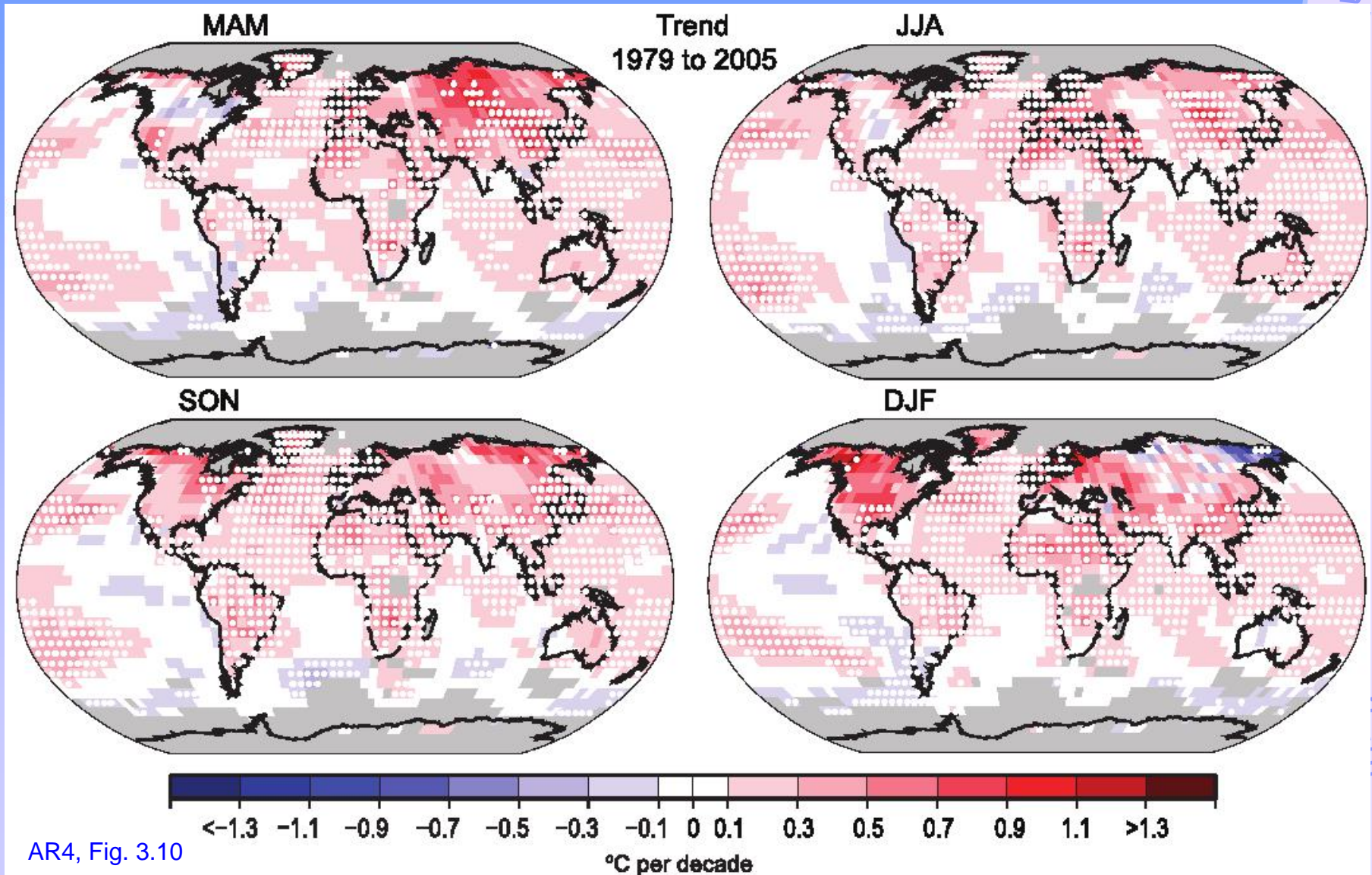
NORTHERN HEMISPHERE TEMPERATURE RECONSTRUCTIONS



The past 157 years



Temperature trends



<http://www.gletscherarchiv.de>



1900



2000



Koninklijk Nederlands Meteorologisch Instituut

1.10.2007

Possible causes

- earth's orbit
- sun (brightness)
- internal variability
- volcanoes
- land-use
- aerosols
- GHGs (CO_2 , CH_4 , N_2O)
- ...

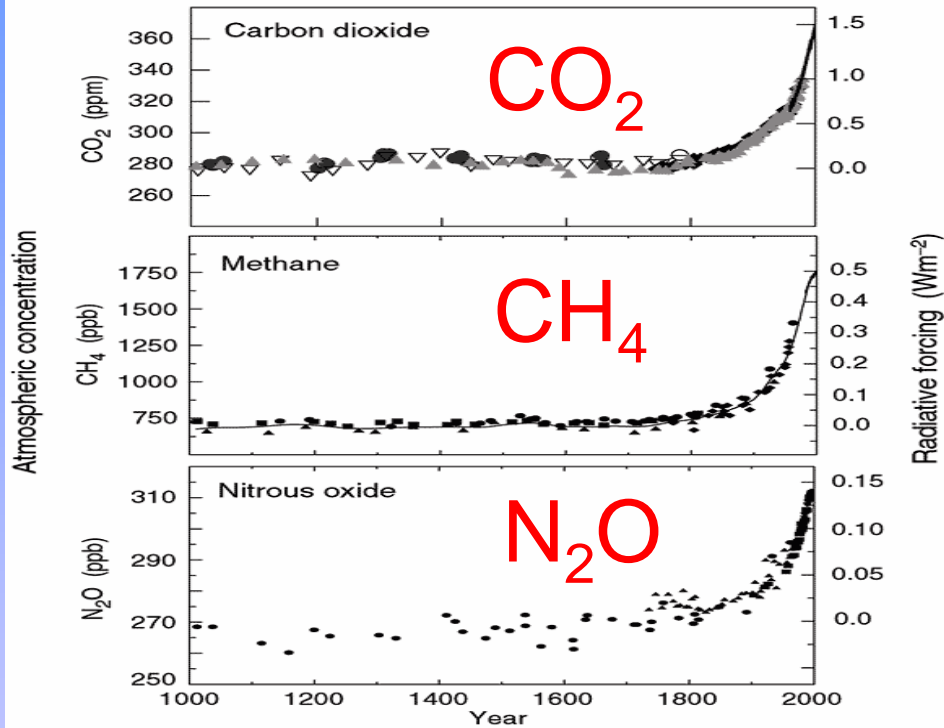
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Increasing GHG concentrations

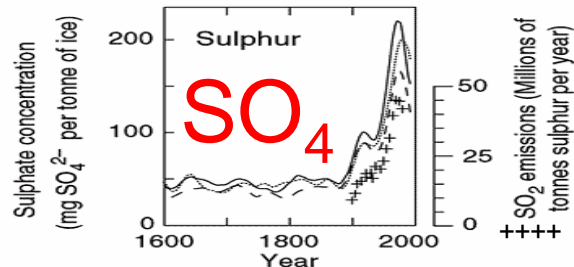


Indicators of the human influence on the atmosphere during the Industrial Era

(a) Global atmospheric concentrations of three well mixed greenhouse gases



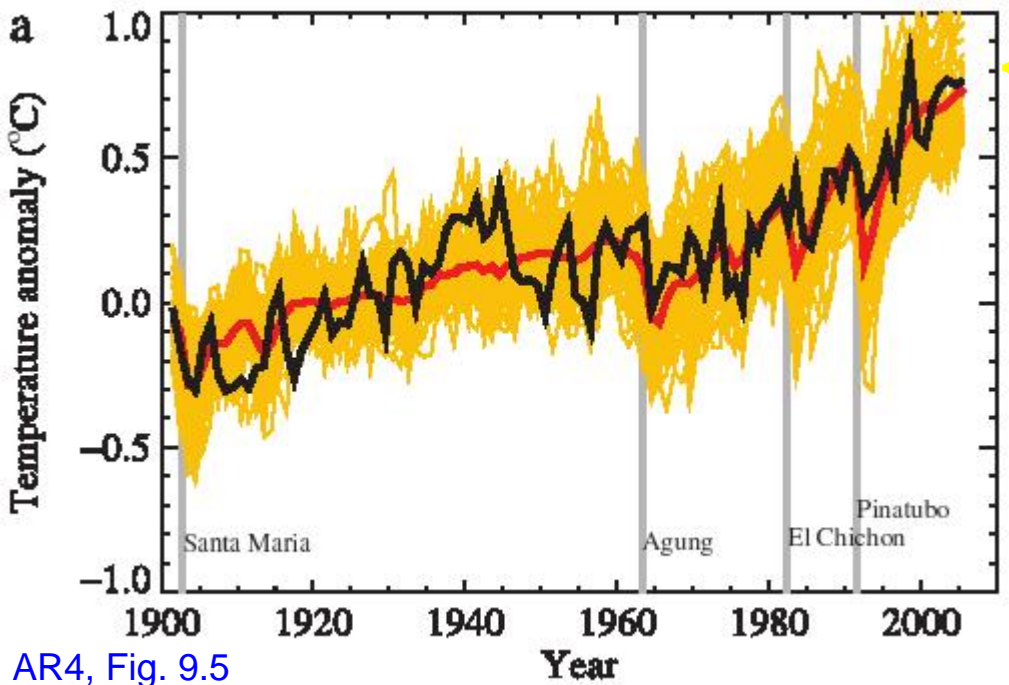
(b) Sulphate aerosols deposited in Greenland ice



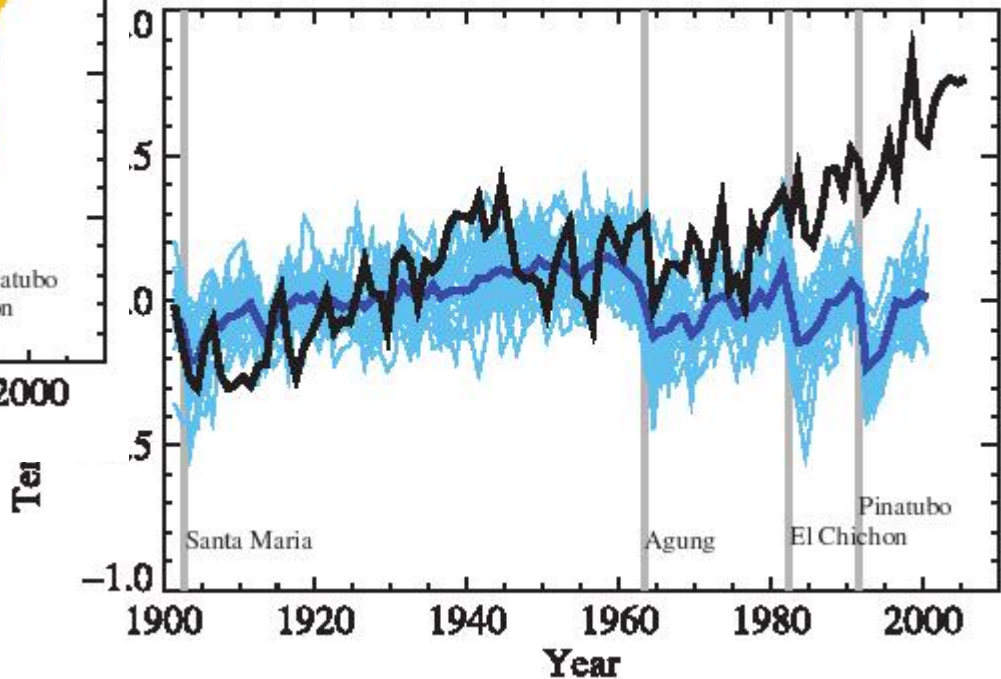
(IPCC, 2001, Fig. SPM-2)

Simulations

← multi model, all forcings



AR4, Fig. 9.5



Multi model, natural forcings =>

Conclusion

The changes after 1950 can only be understood if changing GHG concentrations are taken into account.

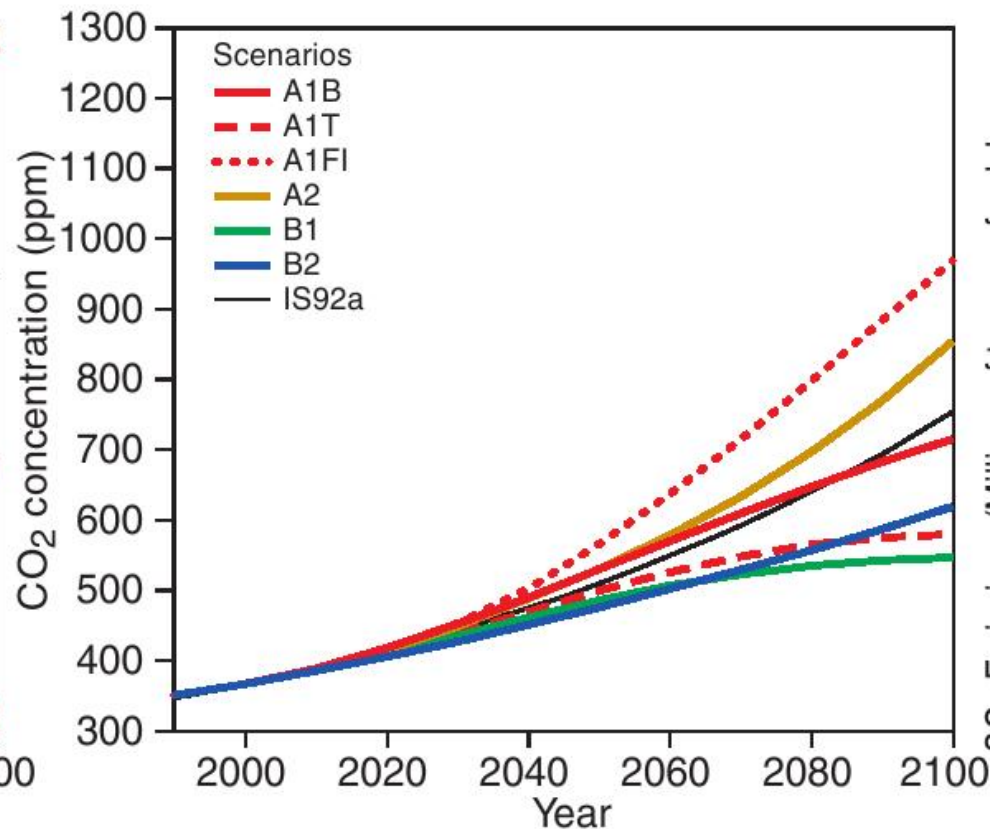
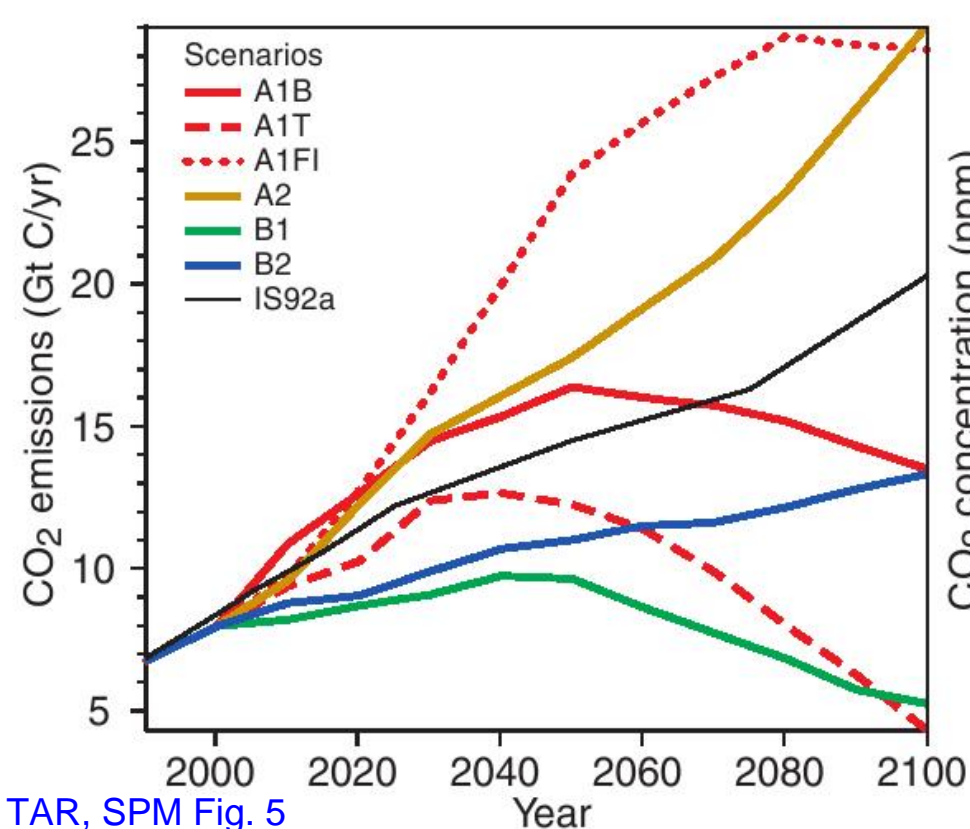
Projections

1. Scenarios
“How will the world change?”
2. Scenarios => CO₂ emissions and concentrations
3. CO₂ concentrations => climate

Emission scenarios

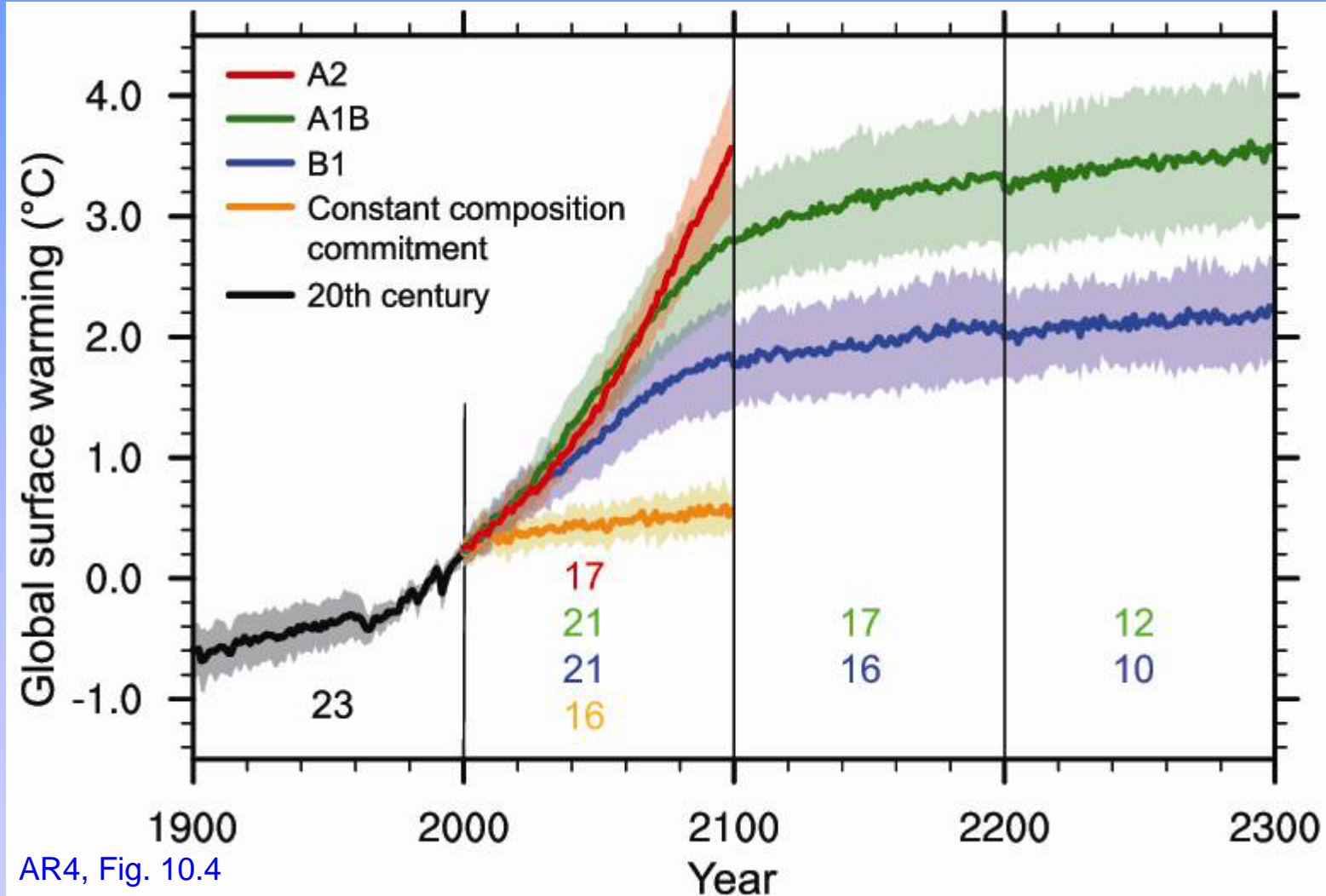
(a) CO₂ emissions

(b) CO₂ concentrations



TAR, SPM Fig. 5

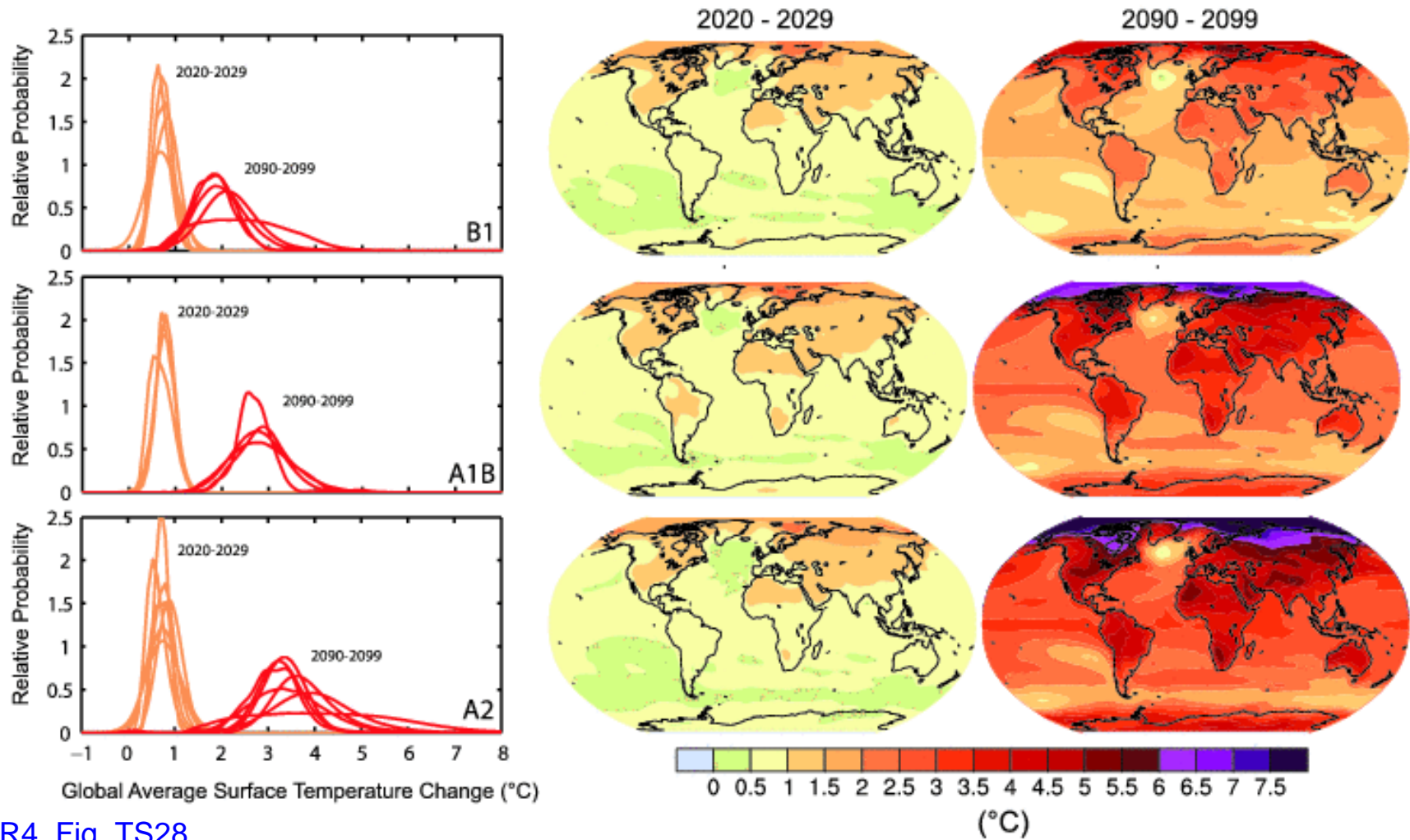
Projected changes



AR4, Fig. 10.4

Projected patterns

PROJECTIONS OF SURFACE TEMPERATURES



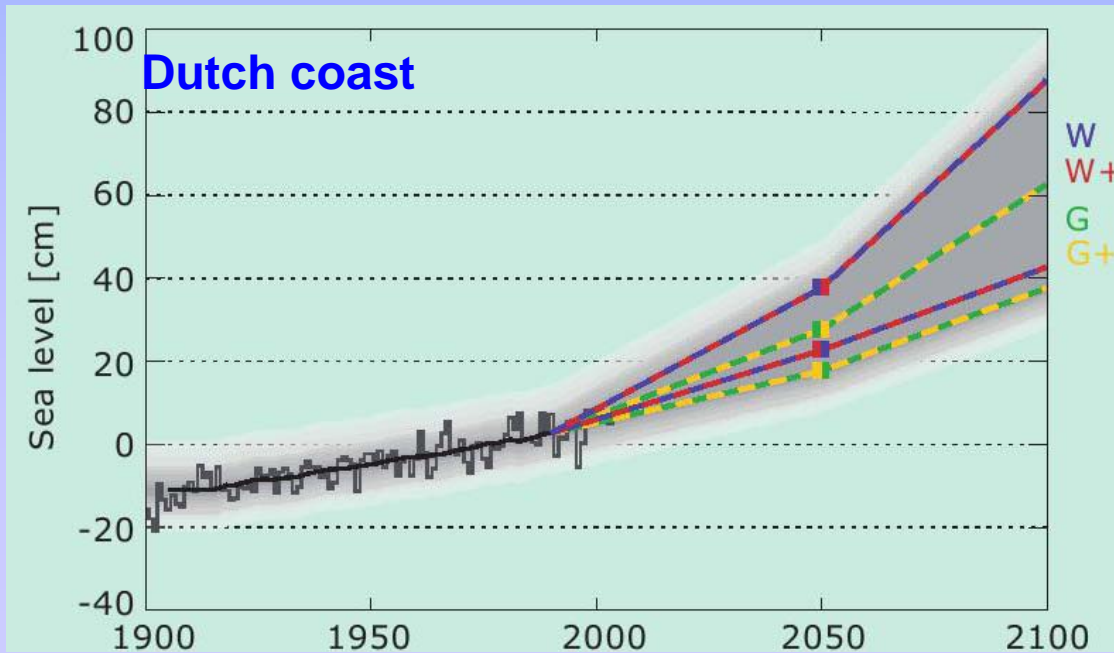


Sea level rise



A1b, 21st century:

- thermal expansion 13-32 cm
- dynamical effects (local) ± 20 cm
- small icecaps 8-15 cm
- Greenland and Antarctica ???



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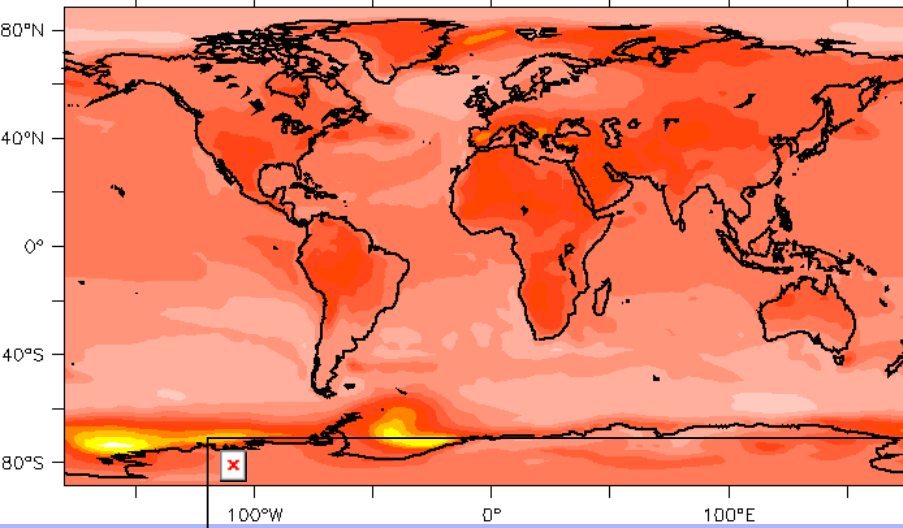
- **17 runs**
- **1950 - 2100**
- **observed GHG's – SRES A1b**
- **ECHAM5/MPI-OM**

Temperature

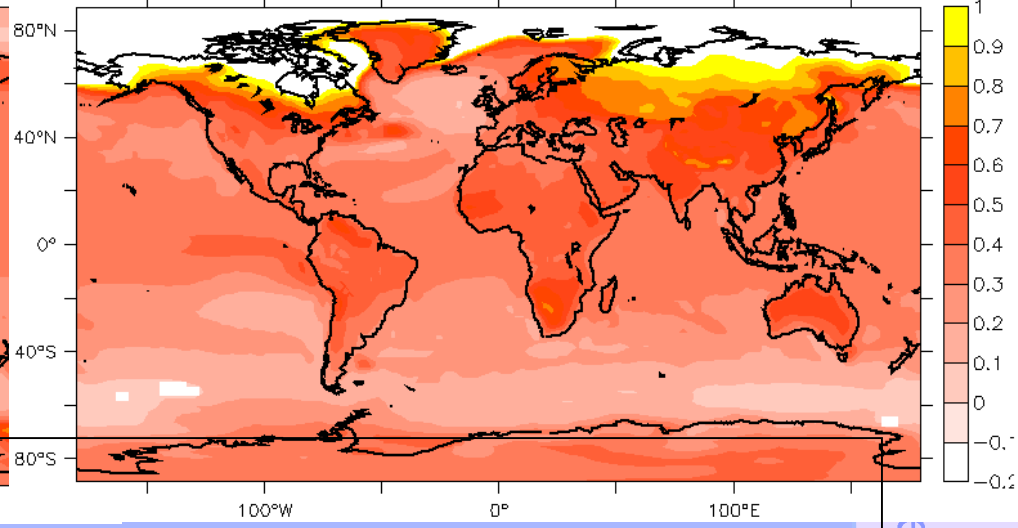


Ko

trend temp2 2001-2100 JJA



trend temp2 2001-2100 DJF

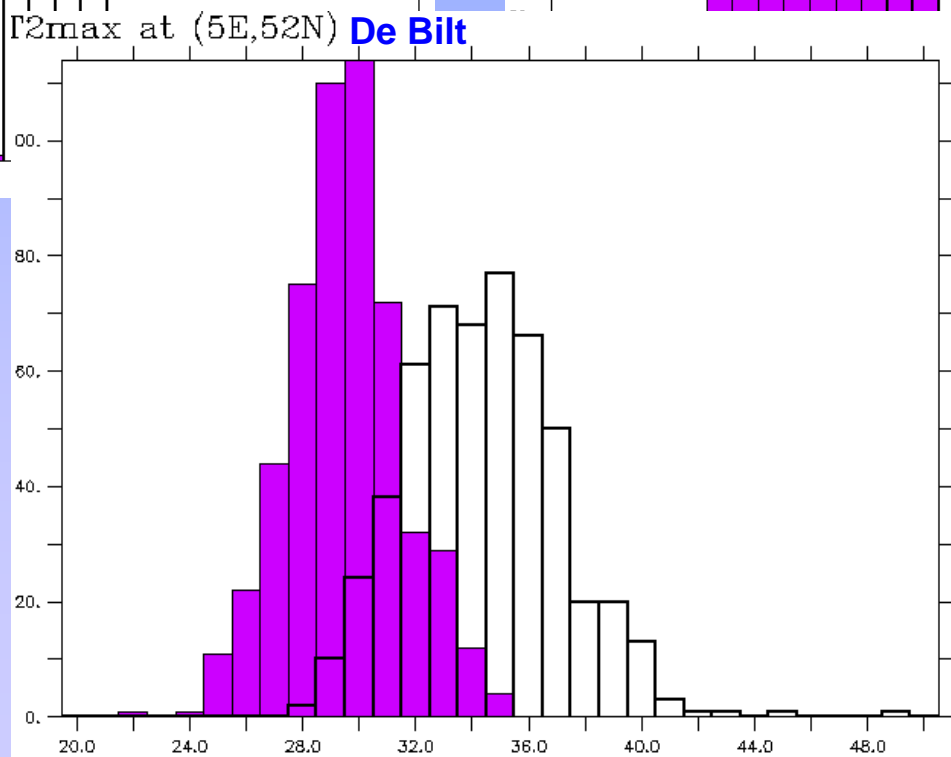
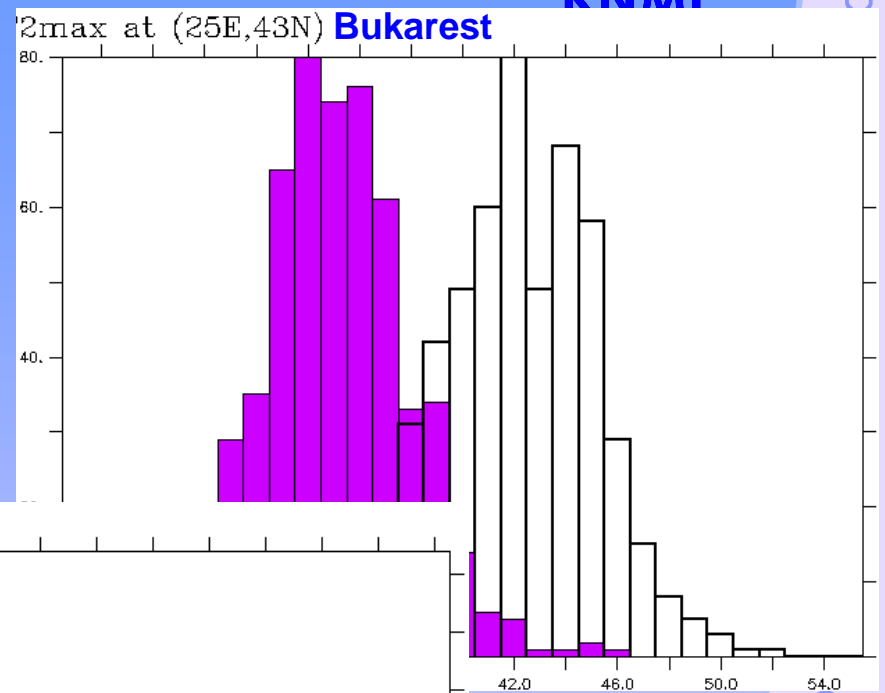
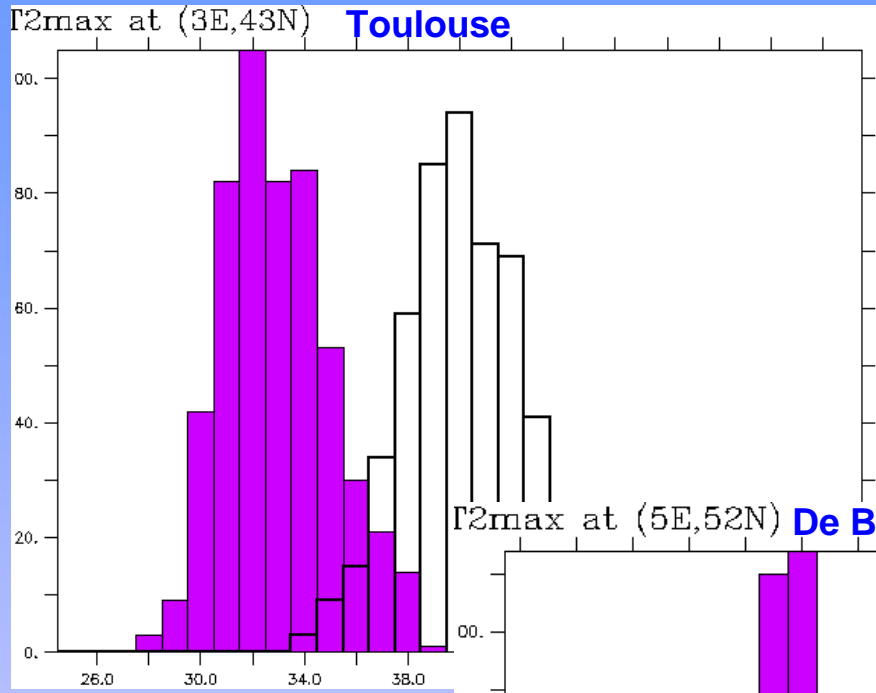


eteorologisch Instituut

Maximum Temperatures



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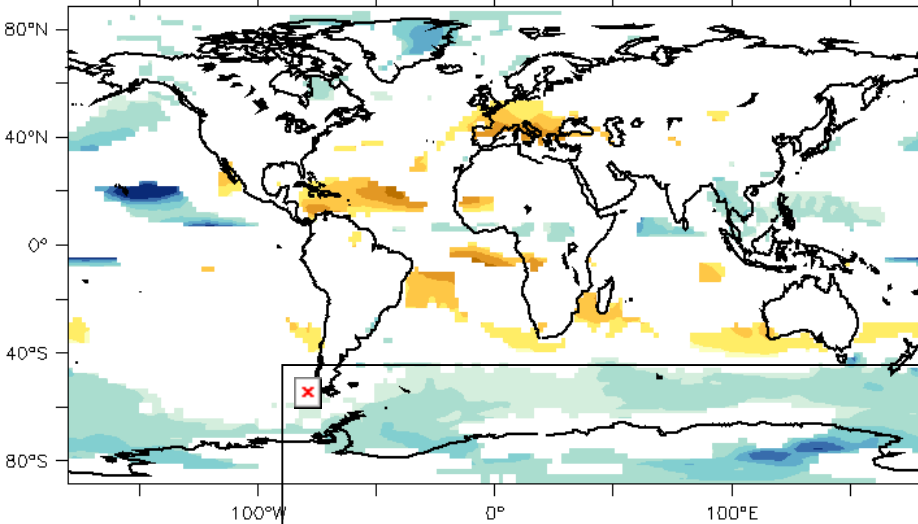
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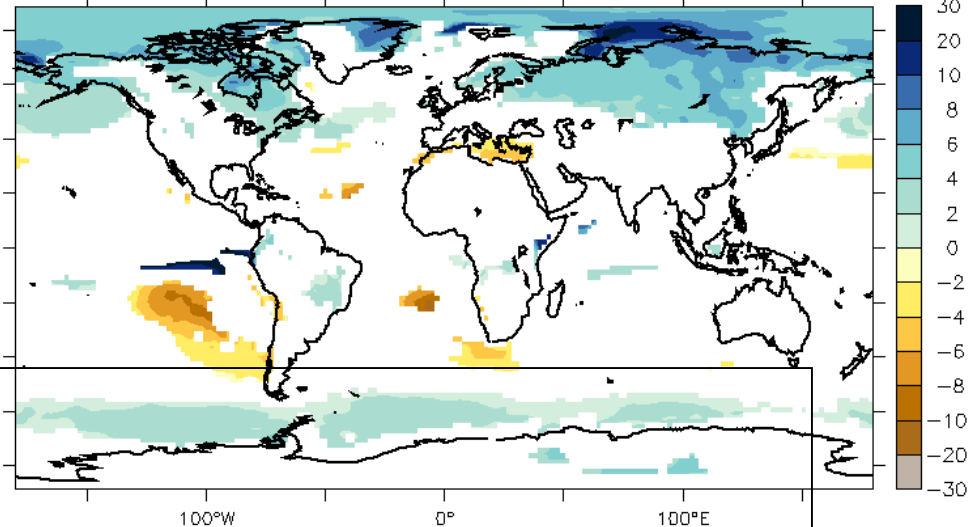


Precipitation

trend precip 2001-2100 JJA

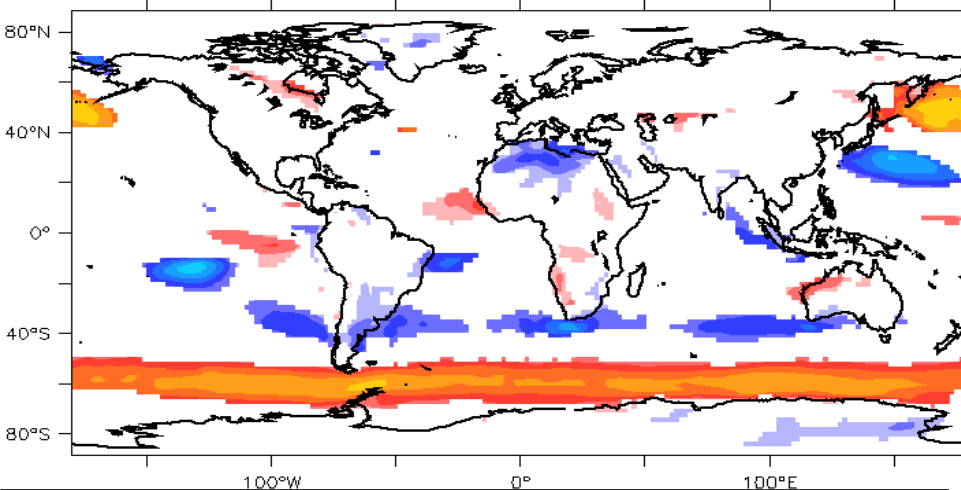


trend precip 2001-2100 DJF

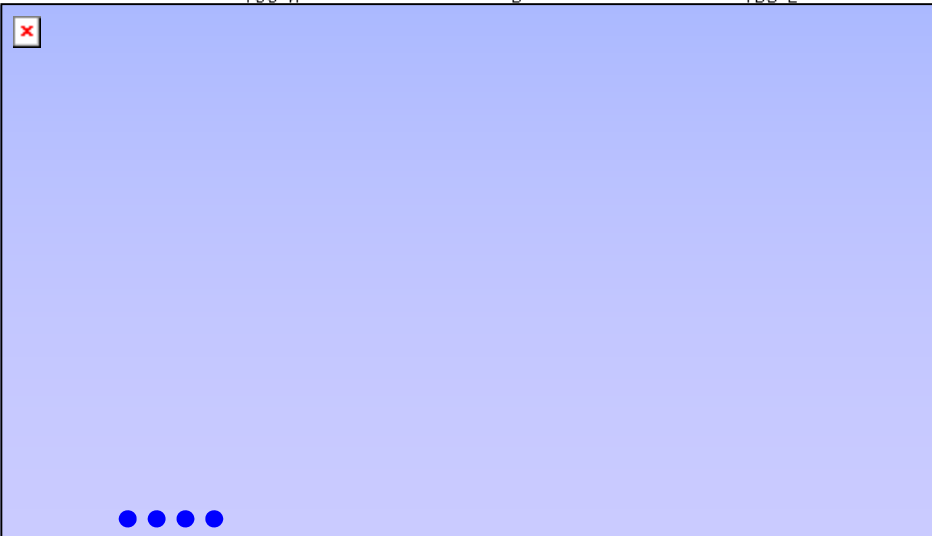
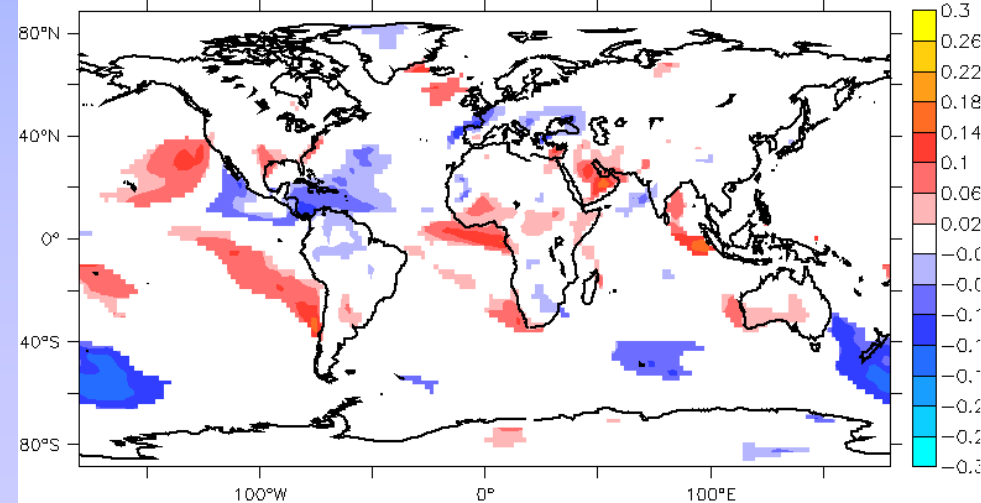


Essence: wind changes

trend u10 2001-2100 DJF



trend v10 2001-2100 JJA

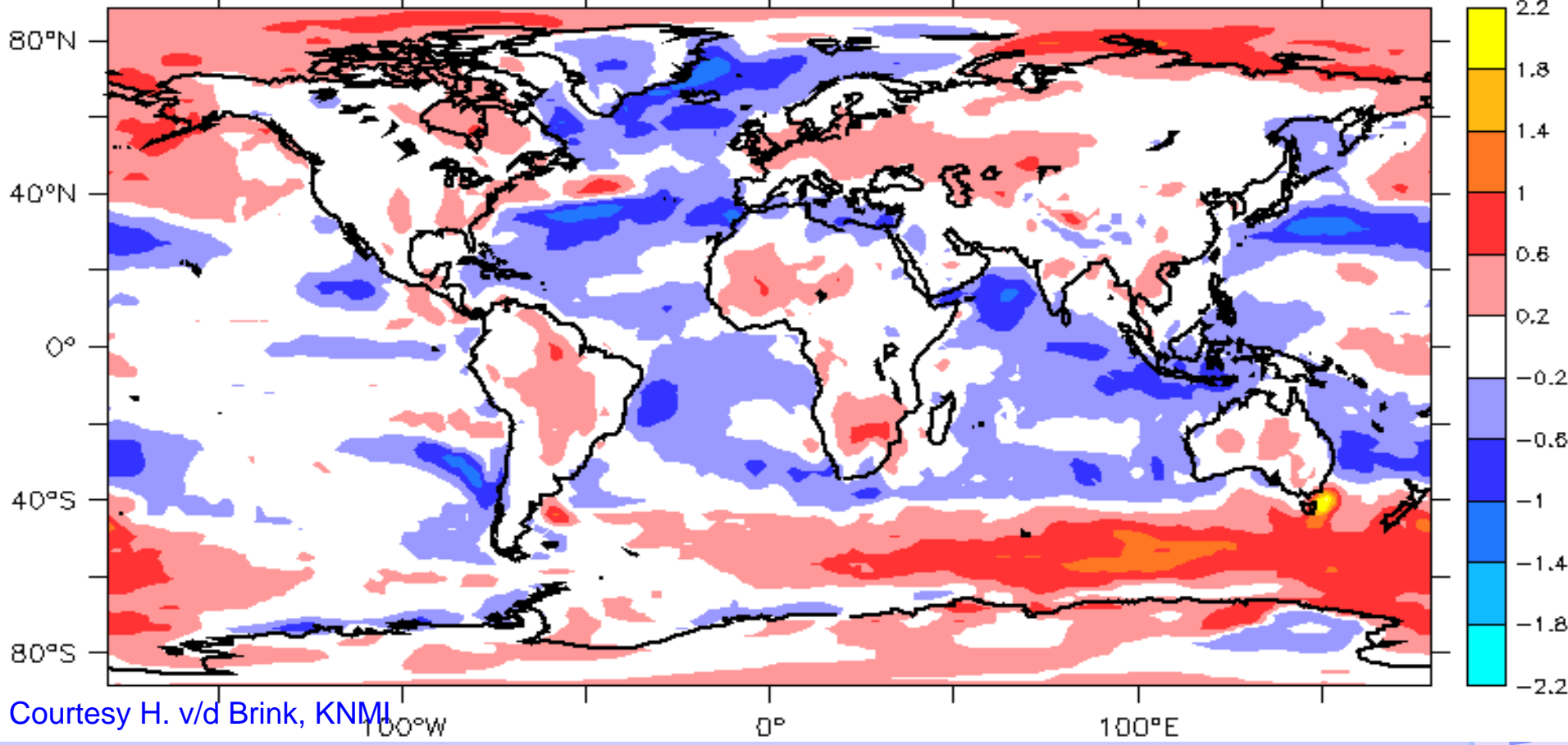




Extreme winds



diff. sqrt(mu), GEV fit to $u \sim 2$



Courtesy H. v/d Brink, KNMI



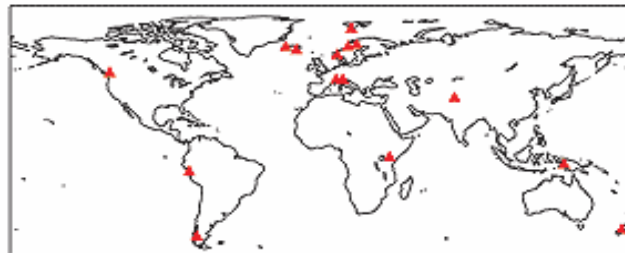
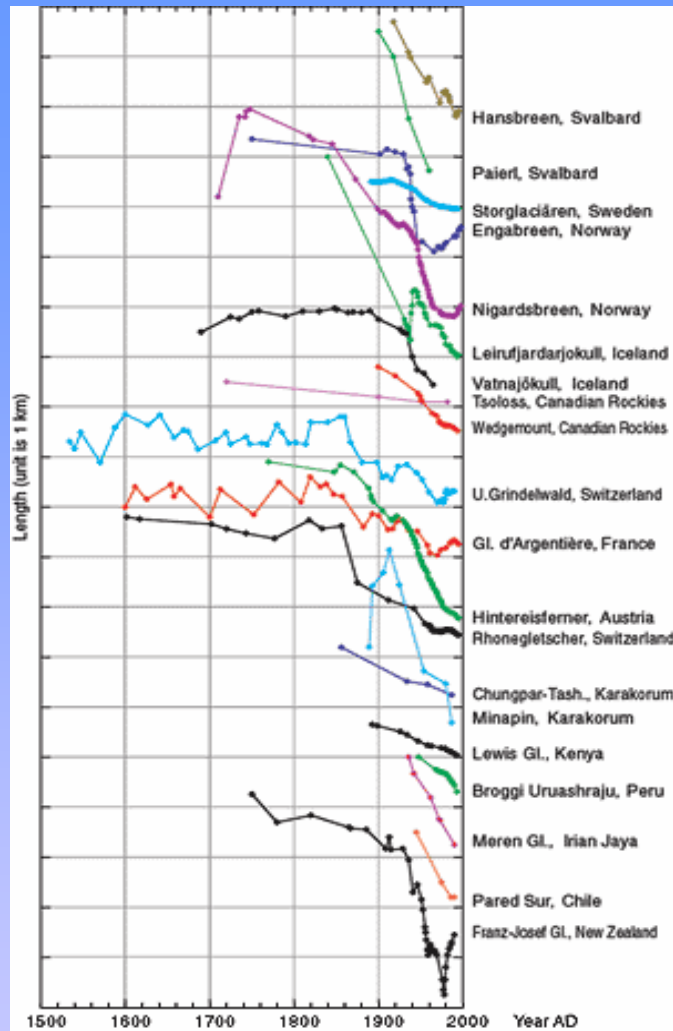


Conclusions

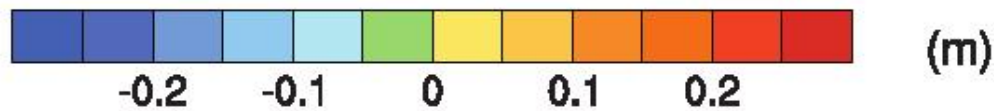
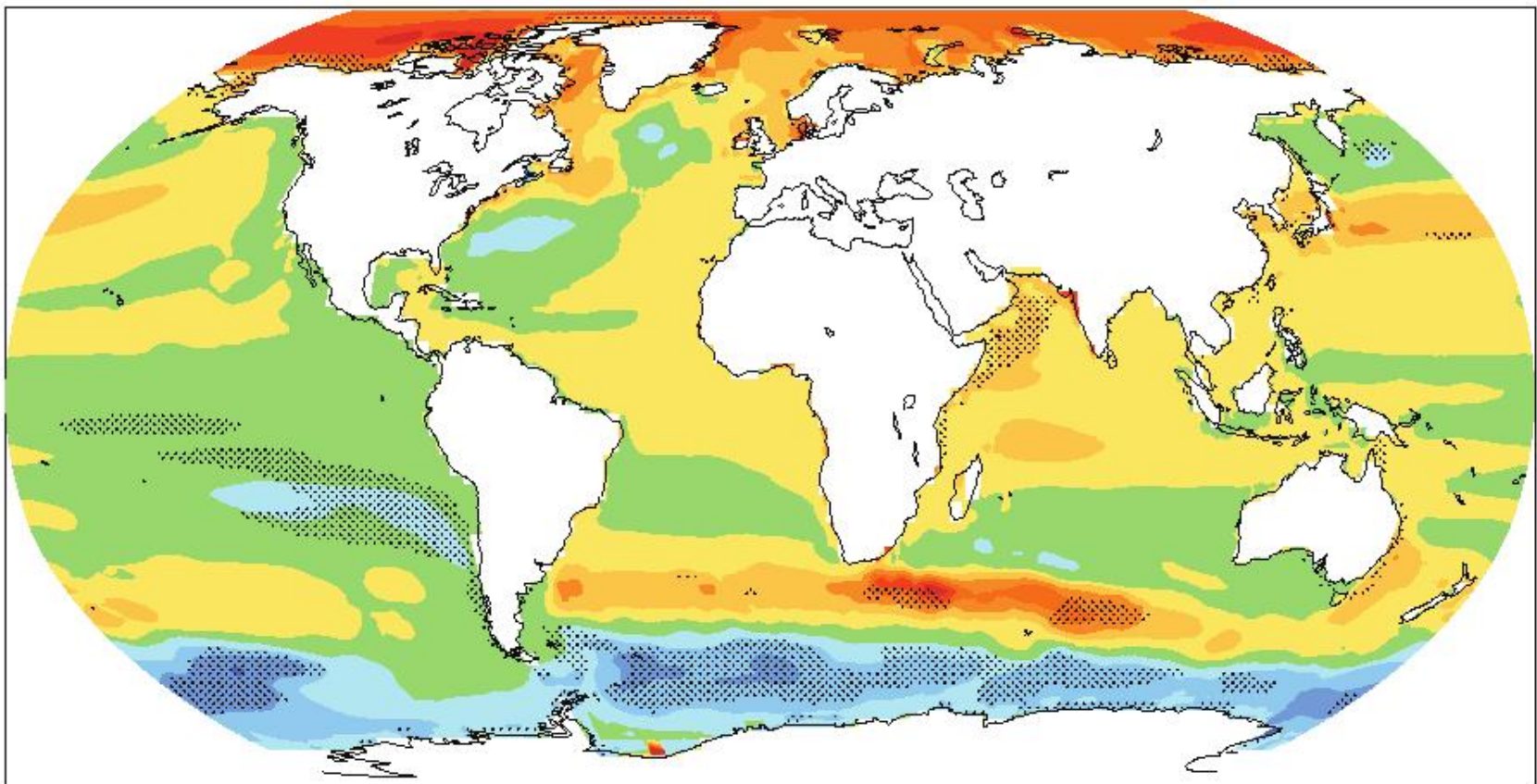
- Climate is already changing and will go on changing:
- Temperatures increase
- Extreme temperatures increase even more
- high latitudes get wetter
- subtropics get drier
- ice is melting
- sea level is rising
- wind regimes move

The End !!!

Glacier lengths



Relative sea level change



Significance of trends - 1

Naïve: z-val:

$$z = \frac{|T_{\text{end}} - T_{\text{begin}}|}{\sqrt{\frac{\sigma_{\text{end}}^2}{17} + \frac{\sigma_{\text{begin}}^2}{17}}} > 1.96$$

=> Every change significant if enough members

Significance of trends - 2

time series: $x(t_i), \quad i = 1, \dots, n_t$

regression line: $\hat{x}(t_i) = a + bt_i$ (ens. average)

regression residuals: $e(t_i) = x(t_i) - \hat{x}(t_i)$

standard error of b : $s_b = \frac{s_e}{\sqrt{\sum(t_i - \bar{t})^2}}$

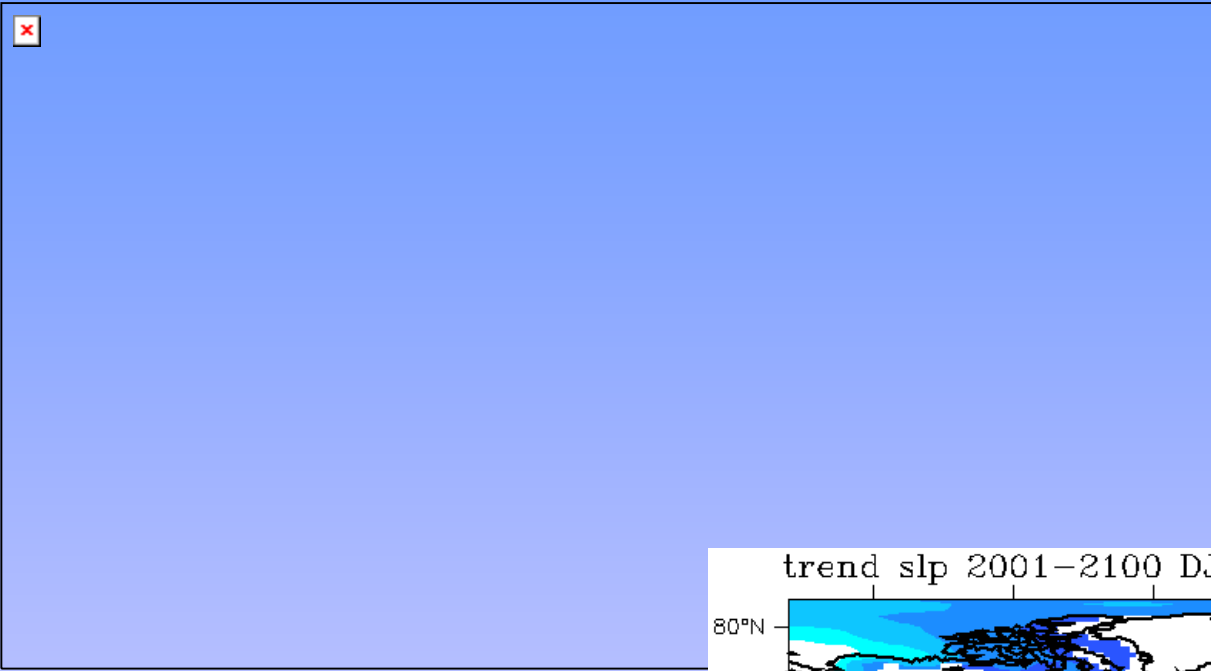
variance of residuals: $s_e^2 = \frac{1}{n_{\text{eff}} - 2} \sum e(t_i)^2$ (ens. members)

t-value: $t_b = \frac{b}{s_b} = b \sqrt{\frac{\sum(t_i - \bar{t})^2}{\sum e(t_i)^2}} (n_{\text{eff}} - 2) > 1.96$

effective dof: $n_{\text{eff}} = n_t \frac{1 - r}{1 + r} \cdot 17$



Sea Level Pressure



trend slp 2001-2100 DJF

